

Modelling unsaturated soil behaviour by using degree of capillary saturation and effective inter-particle stress

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Abstract

This paper discusses the role of the degree of capillary saturation in modelling the coupled hydro-mechanical behaviour of unsaturated soils and proposes a new constitutive model for unsaturated soils by using the degree of capillary saturation and the effective inter-particle stress. In the proposed constitutive model, the shear strength, yield stress and deformation behaviour of unsaturated soils are governed directly by the above two constitutive variables. The model is then validated against a variety of experimental data in the literature, and the results show that a reasonable agreement can be obtained using this new constitutive model.

Keywords: unsaturated soil, capillary water, effective inter-particle stress.

1. INTRODUCTION

The objective of this study is to establish a constitutive model for unsaturated soils by acknowledging that (i) pore water consists of capillary water and adsorbed water, and (ii) they contribute very differently to the constitutive behaviour of unsaturated soils. Because the capillary water exists among soil particles and the pressure of capillary water affects the contact stress among soil particles. Therefore, the stress carried by the capillary water is classified as an inter-particle stress. Compared with the capillary water, the contribution from the adsorbed water to the shear strength and deformation of a soil is very limited (Konrad and Lebeau, 2015; Lu et al., 2010; Zhou et al., 2016). In principle, this is because adsorbed water wraps the surface of each soil particle. The pressure of the adsorbed water should be more likely treated as the internal pressure inside of a particle-surface water system and thus almost does not affect the contact stress among the systems. Thus, the stress carried by the adsorbed water is reasonably treated as an intra-particle stress.

One of the most widely-used constitutive variables for unsaturated soil modelling is the Bishop effective stress (Bishop, 1959), which was extended from Terzaghi's effective stress for saturated soils. The formulation of σ'_{ij} can be written as follows

$$\sigma'_{ij} = \sigma_{ij} + \chi s \delta_{ij} = \bar{\sigma}_{ij} - u_a + \chi(u_a - u_w) \delta_{ij} \quad (1)$$

where σ_{ij} is the net stress with $\sigma_{ij} = \bar{\sigma}_{ij} - u_a$, $\bar{\sigma}_{ij}$ is the total stress, u_a is the pore air pressure, χ is the effective stress parameter, s is the suction ($s = u_a - u_w$), u_w is the pore water pressure and δ_{ij} is the Kronecker delta.

In this study, the degree of capillary saturation (S') is selected as the basic constitutive variable to highlight that only the capillary water affects the strength and deformation of unsaturated soils. Specifically, on one aspect, the degree of capillary saturation is used for the effective stress parameter, i.e., $\chi = S'$, and the effective stress when $\chi = S'$ is referred to as the effective inter-particle stress (σ'_{ij}) to emphasize that the intra-particle stress associated with the adsorbed water pressure has been

ruled out. On the other aspect, the slope of the NCLs is also a function of the degree of capillary saturation, i.e., $\lambda = \lambda(S')$ to underline that the mechanical state of an unsaturated soil is related to the capillary water only. It is important to note that, although the constitutive relationship is initially established in the space of $\{\sigma'_{ij}, S'\}$, it can be generalised in the space of primary variables $\{\sigma'_{ij}, s, S\}$ that is in accordance with variables adopted for finite element methods (Sheng et al., 2003; Zhang and Zhou, 2016; Zhou and Zhang, 2015).

2. CONSTITUTIVE EQUATIONS

2.1. Hydraulic Equations

Sharing the theoretical concept delivered by Or and Tuller (1999) and practical method by Khlosi et al. (2008), a simple equation of water retention curve was proposed to consider capillarity and adsorption separately (Zhou et al., 2016), which can be written as follows:

$$S = S' + S'' \quad (2)$$

where S is the degree of saturation; S' and S'' are the capillary component and adsorptive component of the degree of saturation respectively. Following Khlosi et al. (2008), the two-parameter equation proposed by Kosugi (1996) to quantify the capillary component (i.e., capillary water retention curve, CWRC) was employed here:

$$S' = (1 - S'')C(s), \text{ and } C(s) = \frac{1}{2} \operatorname{erfc}\left((\sqrt{2}\zeta)^{-1} \ln(s/s_m)\right) \quad (3)$$

where $\operatorname{erfc}()$ the complementary error function, s_m the suction that corresponds to the median pore radius (r_m), and ζ^2 the variance of the log-transformed pore radius. s_m varies from s_{mR} corresponding to the main drying branch to s_{mA} corresponding to the main wetting branch. Zhou (2013) stated that the contact angle (θ) for the main drying branch and the main wetting branch are equal to the receding contact angle (θ_R) and advancing contact angle (θ_A), respectively. Also, only on the main drying/wetting branches, the contact angle is independent of the change of the suction. For the scanning processes, the contact angle is approaching to θ_R for drying and to θ_A for wetting. Based on the concept introduced by Zhou (2013), the term $C(s)$ can be revised to consider hysteretic behaviour due to the variation of contact angle caused by the suction change as follows

$$C(s) = \frac{1}{2} \operatorname{erfc}\left((\sqrt{2}\zeta)^{-1} \ln\left[(s \cos \theta_R)/(s_{mR} \cos \theta)\right]\right) \quad (4)$$

where θ_R and θ_A are the receding and advancing contact angles respectively. For simplicity, θ_R is usually assumed to be 0 and θ_A can be calibrated by the main wetting branch. s_{mR} stands for the suction that corresponds to the median pore radius (r_m) in the main drying process, which can be calibrated by the main drying branch. The variation of the contact angle due to suction has been provided by Zhou (2013).

$$d\theta = -\frac{\beta ds}{s \tan \theta} \text{ and } \beta = \begin{cases} \left[(\cos \theta - \cos \theta_R)/(\cos \theta_A - \cos \theta_R)\right]^b & ds \geq 0 \\ \left[(\cos \theta_A - \cos \theta)/(\cos \theta_A - \cos \theta_R)\right]^b & ds < 0 \end{cases} \quad (5)$$

where b is a parameter to adjust the rate of contact angle change due to suction change, which can be calibrated by scanning wetting or drying tests.

In addition to the contact angle hysteresis, the mechanical loading changes soil's pore distribution and further affects its water retention behaviour (Sheng and Zhou, 2011; Zhou and Sheng, 2015; Zhou et al., 2012a, b). The mechanical compression due to net stress increase decreases the median pore radius (r_m) as well as the variance (ζ^2) (Oualmakran et al., 2016). This phenomenon is termed as the mechanical shift of the water retention curve. The following equation is proposed

$$r_m = r_{m0} (1 - \varepsilon_{v\sigma})^a \quad \text{and} \quad \zeta^2 = \zeta_0^2 (1 - \varepsilon_{v\sigma})^a \quad (6)$$

where r_{m0} and ζ_0^2 is the median pore radius and the variance at the reference state (i.e., $\varepsilon_{v\sigma} = 0$), $\varepsilon_{v\sigma}$ is the volumetric strain due to the mechanical loading, a is a parameter to consider the mechanical effects on pore size distribution. Per the Young-Laplace equation, for a given contact angle, suction is in inverse proportion to pore radius. Therefore, term $C(s)$ can be further upgraded to the following expression to consider the mechanical shift.

$$C(s) = \frac{1}{2} \operatorname{erfc} \left(\left[\sqrt{2} \zeta_0 (1 - \varepsilon_{v\sigma})^{a/2} \right]^{-1} \ln \left[\left(s (1 - \varepsilon_{v\sigma})^a \right) / (s_{mR0} \cos \theta) \right] \right) \quad (7)$$

where s_{mR0} is the suction that corresponds to the median pore radius at the reference state ($\varepsilon_{v\sigma} = 0$) and the receding contact angle ($\theta = \theta_R$), ζ_0^2 the variance of the log-transformed pore radius at the reference state ($\varepsilon_{v\sigma} = 0$). Both can be easily calibrated by test results obtained from the drying branch of conventional water retention experiments with a constant net stress. The volumetric strain due to hydraulic loading (ε_{vs}) has been considered when we calibrate s_{mR0} and ζ_0^2 . This is the reason why only the volumetric strain due to mechanical loading ($\varepsilon_{v\sigma}$) is involved in equation (7).

The adsorptive component S'' (i.e., adsorbed water retention curve, AWRC) can be described by the following equation (Zhou et al., 2016):

$$S'' = (1 - P_{cc}) \Theta_a / \Theta_s \quad (8)$$

where Θ_s is the volumetric water content at the fully saturated state, Θ_a is the maximum volumetric water content due to adsorption ignoring capillary condensation due to the mutual influence of adjacent adsorptive water films, and P_{cc} stands for the possibility of capillary condensation ($0 \leq P_{cc} \leq 1$). $P_{cc} = 1$ if soil is fully saturated ($S = S' = 1$) and $P_{cc} = 0$ if capillary water is equal to zero ($S' = 0$). The simplest equation meets the above requirement of P_{cc} can be written as: $P_{cc} = S'$. Specifically, Θ_a can be described by the equation proposed by Campbell and Shiozawa (1992), i.e.,

$$\Theta_a = \Theta_s A(s), \quad \text{and} \quad A(s) = \alpha (1 - \ln s / \ln s_d) \quad (9)$$

where α is the parameter that is related to the maximum degree of saturation due to adsorption (without considering capillary condensation) when the suction is equal to 1 kPa. s_d is the suction at oven dryness. Experimental results have shown that oven dryness generally corresponds to a finite suction of 10^6 kPa. Therefore, the adsorbed water retention curve can be specified as

$$S'' = A(s)(1 - S') \quad (10)$$

Combining equations (2), (3) and (10) yields the following closed-form equations for WRC, CWRC and AWRC:

$$\left\{ \begin{array}{l} \text{WRC: } S = S' + S'' = \frac{C(s) + A(s) - 2C(s)A(s)}{1 - C(s)A(s)} \\ \text{CWRC: } S' = \frac{C(s) - C(s)A(s)}{1 - C(s)A(s)}, \text{ and AWRC: } S'' = \frac{A(s) - C(s)A(s)}{1 - C(s)A(s)} \end{array} \right. \quad (11)$$

2.2. Mechanical equations

It is difficult to construct an effective stress variable that can solely govern both shear strength and volume change behaviour for unsaturated soils. A practical way for unsaturated soil modelling is to achieve this by two steps: (1) construct an effective stress variable that can govern shear strength behaviour only for unsaturated soils, and (2) develop a volume change equation for unsaturated soils using the defined effective stress variable together with another constitutive variable. Realising the inter-particle water (or water bridges or menisci) is composed of the capillary water only while the adsorptive water forms the water film wrapping particles, the effective inter-particle stress equation for the shear strength of unsaturated soil was suggested by Zhou et al. (2016), highlighting only the capillary water contributes to the shear strength under a given suction.

$$\sigma'_{ij} = \sigma_{ij} + S's\delta_{ij} \quad (12)$$

where σ'_{ij} is the effective inter-particle stress, and S' can be determined by equation (11). In the space of deviator stress (q) and mean effective inter-particle stress (p'), the shear strength of unsaturated soil can be written as

$$q = Mp'. \quad (13)$$

The experimental validation for equations (12) and (13) on predicting unsaturated soil strength can be found in Zhou et al. (2016).

Recent research indicates the degree of saturation or the effective degree of saturation can be used as an additional constitutive variable for modelling volume change behaviour. Realising that the degree of capillary saturation contributes much more to unsaturated soil's mechanical behaviour than the degree of adsorptive saturation, the NCLs of unsaturated soils are proposed as below, by adopting the degree of capillary saturation (S') other than the degree of saturation (S) as the key variable.

$$v = N - \lambda \ln p', \quad (14)$$

where p' is the mean effective inter-particle stress defined in equation (12), v the specific volume, N the intercept of the NCL with the v -axis when $\ln p' = 0$, and λ the elastoplastic compression index representing the slope of the NCL, which is assumed as a function of the degree of capillary saturation, i.e. $\lambda = \lambda(S')$.

$$\lambda = \lambda_0 - (\lambda_0 - \lambda_d)(1 - S')^k, \quad (15)$$

where λ_0 is the elastoplastic compression index for the fully saturated soil, λ_d is the elastoplastic compression index for when the soil contains zero capillary water (i.e. oven dryness), and k is a coupling parameter which can be determined by a drying test. For the elastic response, the following equation is employed.

$$dv = -\kappa dp' / p', \quad (16)$$

where κ is the elastic compression index representing the slope of the unloading and reloading line (URL). For κ , an equation like λ is employed.

$$\kappa = \kappa_0 - (\kappa_0 - \kappa_d)(1 - S')^k, \quad (17)$$

where κ_0 is the elastoplastic compression index for the fully saturated soil, and κ_d is the elastoplastic compression index for when the soil contains zero capillary water (i.e., oven dried). For the clayey soil, the capillary water can only be fully removed by very high suction and the compressibility of the clayey soil at a very high suction is far less than its compressibility at the fully saturated state. Therefore, for simplicity, in this paper, we assume $\lambda_d = \kappa_d = 0$ to simplify equation (15) and (17) as

$$\kappa = \kappa_0 - \kappa_0(1 - S')^k, \text{ and } \lambda = \lambda_0 - \lambda_0(1 - S')^k. \quad (18)$$

The yield surface (or the loading collapse surface) function in the isotropic stress states

$$f_{LC} = (p')^{1-(1-S')^k} - p'_c = 0. \quad (19)$$

The Modified Cam-clay model is employed here to extend the isotropic yield surface to a triaxial stress state

$$f = p' + \frac{q^2}{M^2 p'} - (p'_c)^\delta = 0, \text{ where } \delta = \frac{1}{1-(1-S')^k}. \quad (20)$$

3. VALIDATIONS

Cunningham et al. (2003) presented the results of a series of isotropic/triaxial compression tests in both CS and CW conditions on a reconstituted silty clay that comprised a mix of 20% pure Speswhite kaolin, 10% London clay and 70% silica silt. The slurry soil was isotropically preconsolidated to 130 kPa, i.e. the initial yield stress is 130 kPa. The model parameters for simulations are listed in Table 1.

Table 1. Model parameters for the reconstituted unsaturated kaolin

Mechanical parameters (4)				Hydraulic parameters (6)					
λ_0 (-)	κ_0 (-)	M (-)	k (-)	s_{mR0} (kPa)	ζ_0 (-)	α	θ_A (°)	a (-)	b (-)
0.044	0.007	1.25	3.0	1300	0.7	0.4	50	2.6	1.2

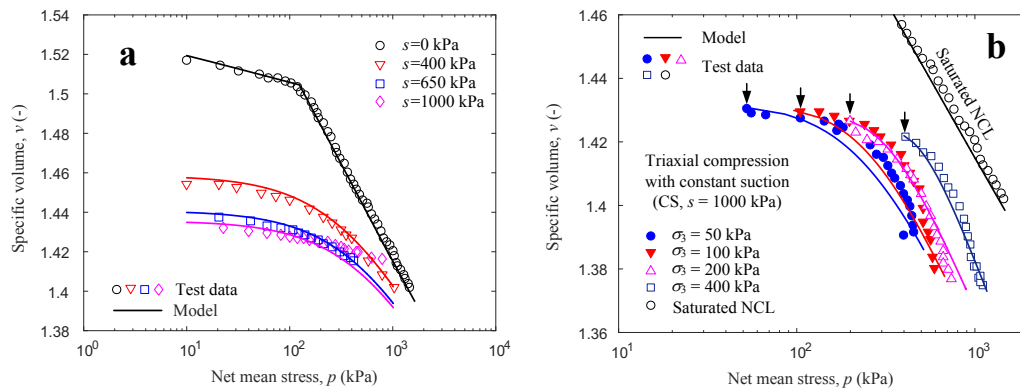


Figure 1. Volume change of the reconstituted kaolin and model simulations

Figure 1a shows the volume changes for the reconstituted kaolin dried to different suctions (0, 400, 650 and 1000 kPa) due to mechanical loading (net mean stress) in the isotropic stress state. The comparison between the test data and the model prediction indicates the model capacity in mimicking the volume change behaviour of the reconstituted kaolin with various suctions due to mechanical loading. Figure 1b shows the volume change against the net mean stress for four triaxial compression tests, which are conducted under a constant suction of 1000 kPa but under different confining stresses (50, 100, 200 and 400 kPa). The proposed model is employed to predict the relationship between the specific volume and the net mean stress observed in four different triaxial compression tests (see Figure 1b). In general, the prediction matches the observation reasonably well.

4. CONCLUSIONS

A new constitutive model for unsaturated soils using the degree of capillary saturation and effective inter-particle stress is proposed in this paper, where the shear strength, yield stress and deformation behaviour of unsaturated soils are governed by two constitutive variables. The proposed constitutive model can capture the observed mechanical and hydraulic behaviours with a limited number of parameters. The capacity of the proposed model has been validated against a variety of experimental data in the literature.

REFERENCES

- Bishop, A.W., (1959) The principle of effective stress. *Teknisk Ukeblad* 106, 859-863.
- Campbell, G.S., Shiozawa, S., 1992. Prediction of hydraulic properties of soils using particle-size distribution and bulk density data., in: van Genuchten, M.T., Leij, F.J., Lund, L.J. (Eds.), *Indirect methods for estimating the hydraulic properties of unsaturated soils*, Univ. of California, Riverside, pp. 317-328.
- Cunningham, M.R., Ridley, A.M., Dineen, K., Burland, J.B., (2003). *Geotechnique* 53, 183-194.
- Khlosi, M., Cornelis, W.M., Douaik, A., van Genuchten, M.T., Gabriels, D., (2008). *Vadose Zone Journal* 7, 87-96.
- Konrad, J.-M., Lebeau, M., (2015). *Canadian Geotechnical Journal* 52, 2067-2076.
- Kosugi, K., (1996). *Water Resources Research* 32, 2697-2703.
- Lu, N., Godt, J.W., Wu, D.T., (2010). *Water Resources Research* 46, W05515.
- Or, D., Tuller, M., (1999). *Water Resour. Res.* 35, 3591-3605.
- Oualmakran, M., Mercatoris, B.C.N., François, B., (2016). *Canadian Geotechnical Journal* 53, 1902-1909.
- Sheng, D., Sloan, S.W., Gens, A., Smith, D.W., (2003). *International Journal for Numerical and Analytical Methods in Geomechanics* 27, 745-765.
- Sheng, D., Zhou, A.N., (2011). *Canadian Geotechnical Journal* 48, 826-840.
- Zhang, Y., Zhou, A., (2016). *International Journal for Numerical and Analytical Methods in Geomechanics* 40, 2353-2382.
- Zhou, A., Huang, R.-Q., Sheng, D., (2016). *Canadian Geotechnical Journal* 53, 974-987.
- Zhou, A.N., (2013). *Computers and Geotechnics* 49, 36-42.
- Zhou, A.N., Sheng, D., (2015). *Computers and Geotechnics* 63, 46-66.
- Zhou, A.N., Sheng, D., Sloan, S.W., Gens, A., (2012a). *Computers and Geotechnics* 43, 178-187.
- Zhou, A.N., Sheng, D., Sloan, S.W., Gens, A., (2012b). *Computers and Geotechnics* 43, 111-123.
- Zhou, A.N., Zhang, Y., (2015). *Computational Mechanics* 55, 943-961.